Radar Remote Sensing of the Meteo-Marine Parameters in the Baltic Sea

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Motivation

- To develop/improve, validate and demonstrate the value of meteomarine parameters derived from different radar data
 - Assess current state-of-the-art methods in estimating meteo-marine parameters
 - High resolution TerraSAR-X/TanDEM-X StripMap imagery
 - Medium resolution Sentinel-1A/B IW swath imagery
 - Marine radar imagery
 - Validate wave retrieval methods in the Baltic Sea
 - XWAVE_C
 - Pleskachevsky et al. 2016, ISPRS, 119; Rikka et al. 2018, IJRS, 39(4)
 - CWAVE_S1-IW
 - Pleskachevsky et al. (submitted to IJRS); Rikka et al. 2018, Remote Sensing, 10(5)
 - Method for marine radar
 - Rikka et al. (submitted to IEEE Geoscience and Remote Sensing Letters)
 - Validate wind retrieval methods (XMOD-2 and CMOD)
 - Rikka et al. 2018, IJRS, 39(4); Rikka et al. 2018, Remote Sensing, 10(5)
- to compare different radar (TS-X, Sentinel-1, marine radar) wave retrievals with (operational) wave model results
- to examine the added benefits of radar data to maritime situation awareness in the Baltic Sea



Meteo-marine climate in the Baltic Sea

- Complex coastline
- Thousands of islands
- Dominant wind direction
 - Sector 180° 315° (S NW)
 - Frequently observed slanting fetch cases, up to 50°
- Dominant wave period
 - 2-8 s
 - Small swell component in H_s
- Dominant wave height
 - H_s between 0–3 m
 - Up to about 10 m observed
 - Dependent of the region
 - Clear annual cycle
- Short wave "memory"
- Hardly recognizable wave pattern on SAR imagery
- Noisy SAR images



Leppäranta and Myrberg, 2009

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Radar imaging of sea surface: SAR







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Radar imaging of sea surface: marine radar

- Marine radar imaging introduce additional effects
 - Very high incidnce angles
 - Shadowing
 - Scattering from micro breakers, i.e. whitecapping





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Data - in situ measurements, radar, wave model

Dev. – algorithm developement; Val. – validation; Comp. – comparison with *in situ* or wave model results; Stat. – seasonal or regional statistics

Sensor	Radar	Pixel size	Temp.	Spatial	Period	No.	Purpos	In situ	Wave
	wavelength		res.	coverage		of	е	collocation	model
						images		S	collocations
TS-X	X-band	3×3 m	On	30× up			Πον	117 LI	
TD-X	3.1 cm		demand	to about	2012-	02	Dev.	102//	55 <i>L</i> _P , γ _P SWAN
				250 km	2017	92	Val.	102 U ₁₀	
							Comp.	44 L _P , γ _P	
Sentin	C-band	10×10 m	1 – 2	250× up		15	Val	52 H _s	49314 H _s
el-1 IW	5.5 cm		days	to about	2015-		VdI.	358 U ₁₀	WAM
				few 10 ³ km	2016	460	Comp.	101 <i>H</i> _s	201 <i>H</i> _s
						400	Stat.		
Marin	X-band	5×5 m	1 h	About		550	Πον	1678 H	-
e radar	3.2 cm			10 km	10.10		DEV.		
				from radar	10.10	Jan.	Val.	1464 H _s	-
				tower	14.11.10	& Jun.			
						2017			

 $H_{\rm S}$ total significant wave height

 U_{10} wind speed

 $L_{\rm P}$ peak wave lenght

 $\gamma_{\rm P}$ peak wave propagation direction

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Data – *in situ* measurements, radar, model Wave



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SAR methods: wind

- Sea state is strongly dependent on local wind characteristics
- XMOD-2 and CMOD
- $\sigma_0(U,\theta,\phi) = B_0^p(U_{10},\theta)(1+B_1(U_{10},\theta)\cos(\phi)+B_2(U_{10},\theta)\cos(2\phi))$
 - σ_0 Normalized Radar Cross Section (NRCS)
 - U_{10} wind speed
 - ϕ wind direction relative to flight direction
 - θ local incidence angle
 - B_i tuned function parameters for XMOD-2 and CMOD separately
 - With polarisation ratio for XMOD-2 (Li and Lehner, 2014):
 - $PR = \frac{\sigma_0^{VV}}{\sigma_0^{HH}} = X_0 EXP(X_1\theta)$, where X₀ and X₁ are tuning coefficients
- According to Monaldo et al. 2016, separate GMFs are used to receive wind speed
 - CMOD4 with Thompson, D. R., et al. (1998) PR for HH polarization and CMOD5.N for VV polarization
- Wind direction from Weather Research and Forecasting Model (WRF) is used (Skamarock et al. 2005)
- WRF wind direction are interpolated to the sea state calculation grid

Radar methods: sea state

- Calculation of NRCS from pixel's digital number
- Artefacts filtering
- Subscene normalization $\sigma_n(x, y) = \frac{\sigma_0(x, y) - \langle \sigma_0 \rangle}{\langle \sigma_0 \rangle}$
- Fast Fourier Transform
- Empirical function without transformation into wave spectra
- Additional Grey Level Cooccurance Matrix (GLCM) image statistics
- General methods are based on validation data matchups from open source measurement data from all over the



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Sea state parameter estimation



Quality Control: Buoys (location) and Wave model results (spatial distribution)

- XWAVE_C Pleskachevsky et al. 2016, *ISPRS*, 119
- CWAVE_S1-IW Pleskachevsky et al. (*submitted to IJRS*))

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Radar methods: sea state

- Energy of image spectrum retrieved by FFT operaator
 - $E_{IS} = \int_{k_x^{min}}^{k_x^{max}} \int_{k_y^{min}}^{k_y^{max}} IS(k_x, k_y) dk_x dk_y$ • $k = \sqrt{k_x^2 + k_y^2}$ where k_{max} and k_{min} depend on radar data used
- Significant wave height
 - $H_{\rm S} = a_0 \sqrt{B_0 E_{IS} \tan(\theta)} + \sum_{i=1}^n a_i B_i$
 - θ is local incidence angle
 - a_i are coefficients, and B_i are functions of spectral parameters, wind and GLCM results depending on data/sensor
- Empirical algorithm for marine radar data
 - *H*_s estimation based on image spectra *E*_{*IS*}
 - Calculated parameters are tested against measured in situ values
 - Best-fit trendline technique
 - Pearson correlation coefficient
 - Minimize RMSE
 - $B_0 = f(d, \theta)$
 - $B_1 = f(d, \theta, \overline{x})$, where $\overline{x} = \sum_{i=0}^{2G-2} i P_{x+y}(i)$
 - $B_2 = f(d, \theta, \sigma^2)$, where $\sigma^2 = \sum_{i=0}^{G-1} \sum_{j=0}^{G-1} (i \mu)^2 P(i, j)$
 - *d* distance from radar tower
 - \overline{x} GLCM mean
 - σ^2 GLCM variance
 - P number of collocations in GLCM levels G



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Results: validation

- High agreements between in situ wind and radar-derived wind speed, especially for Sentinel-1 results where RMSE less than 1.5 m s⁻¹
- Radar derived H_s accurate, r slightly less than 0.90, RMSE less than 0.5 m
- High agreement in range of 0 – 3 m (typical for Baltic Sea) between SAR and WAM



 Collocation	TS-X	TS-X	Sentinel-1	Sentinel-	Sentinel-	Marine
pair	TD-X vs. in	TD-X vs. <i>in</i>	vs. in situ	1 vs. in situ	1 vs. WAM	radar vs. <i>in</i>
	situ	situ				situ
Parameter	H _s	<i>U</i> ₁₀	H _s	<i>U</i> ₁₀	H _s	H _s
 r	0.88	0.90	0.88	0.91	0.86	0.78
RMSE	0.32	2.02	0.40	1.43	0.47	0.23
SI	0.33	0.24	0.37	0.19	0.33	0.41
п	117	102	52	357	49314	1678



Sentinel-1 data for regional studies

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atitude

- Wave height up to 7.5 m
- General agreement in the wave height values and location of maximum
- Storm peak area smaller from SAR data
- Storm does not spread as • much to the north as on WAM field
- Maximum H_s higher with SAR-derived results
- Wave field variability (STD) many times larger for SAR dataset
- Variability in wave model fields lost mostly due to wind forcing, local fine-scale wind field variations and gusts are not included in wave model forcing



Rikka et al. 2018, Remote Sensing, 10(5)



Local variability of sea state from high resolution SAR imagery

- General agreement between SWAN wave model results and SAR-derived H_s values
- Wave height, wavelenght and wave propagation direction shows more variability from radar-derived results



Sentinel-1 data for opera

- An independent time series from 1st August 2016 until the end of 2016
- Case 1 mismatch in wave height value on WAM
 - better detailed spatial variability
- Case 3 similar to Case 1 but with more uniform wave field
- Case 2 missing in situ or model data can be covered by SAR
 - **Technical** issues
 - Maintenance of measurement device
- Boos measurement station Södra Östersjön – no *in situ* data since 2011, although common high sea



Coastal radar data for operational service

- Average H_s field during 26.03 – 28.03.2017 for NW storm conditions
- Time series of *in situ* measurements and radar-derive H_s show good agreement during the storm
- Similar H_s field has been shown by other authors for comparable conditions
- Waves propagating into Tallinn Bay between the mainland and Naissaare
- Maximum H_s around the tip of the Paljassaare peninsula
 - Depth around 30 m



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Statistical mapping of coastal wave field

E.g. Kudryavtseva and Soomere (2017) analysed altimetry data over the Baltic Sea

- Data between 1993-2015
- Output resolution about 0.2×0.1°



- Data between 2015-2016 [•] ^g
 Output resolution 3 nm
- Output resolution 3 nm (interpolated to 1 nm grid)



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Conclusions

- Methods to estimate total significant wave heights were improved/developed and validated for the Baltic Sea wave climate conditions
 - XWAVE_C
 for H_s (r = 0.88, RMSE = 0.32 m)

 CWAVE_S1-IW
 for H_s (r = 0.88, RMSE = 0.40 m)

 Marine radar
 for H_s (r = 0.78, RMSE = 0.23 m)

 XMOD-2
 for U_{10} $(r = 0.90, RMSE = 2.02 m s^{-1})$

 CMOD
 for U_{10} $(r = 0.91, RMSE = 1.43 m s^{-1})$
- The statistics show that radar-derived results are suitable for routine monitoring of meteo-marine parameters in the Baltic Sea
- SAR-derived values of geophysical parameters are spatially more variable and would provide more detailed wave field compared to wave model
- SAR-derived results could be used for wave model validation
 - Wind data from SAR could be used as wave model forcing



Conclusions

- SAR-derived wave height and wind speed results can replace measurements or wave model results in poorly sampled areas or in cases when data is missing
- SAR enables to observe coastal wave field variations in the Baltic Sea in more detail compare to other EO sensors (altimetry)
- SAR data enable to perform wave climate studies in seasonal and regional scale
- Based on Paljassaare marine radar data analysis wave height can be monitored with high accuracy in space and time
- Considering all above the radar based wind and wave data would be beneficial for maritime situation awareness applications and routine monitoring/forecasting in the Baltic Sea



Thank you for your attention!



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